

Faster Tropical Upper Stratospheric Upwelling Drives Changes in Ozone Chemistry

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**BOTTOM LINE: Upper Strat Composition Change →
*O₃ Loss Changes → O₃ trends from 2005-2021***

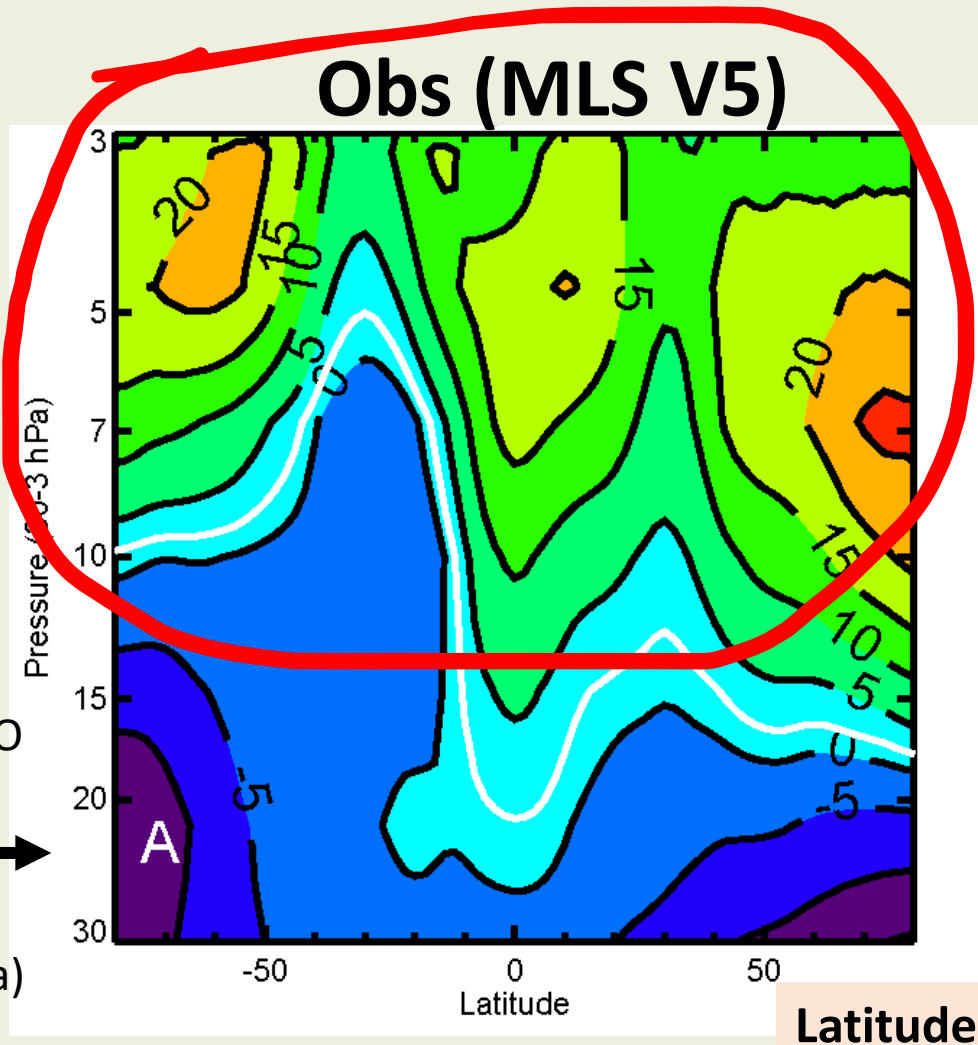
Quasi-Decadal (QD) N₂O Change over 2005-2021

(Percent change of 2013-2021 mean - 2005-2013 mean)

N₂O surface growth over this period = 2%

No drift in MLS N_2O
22 hPa and above

Known negative
drift below 22 hPa)



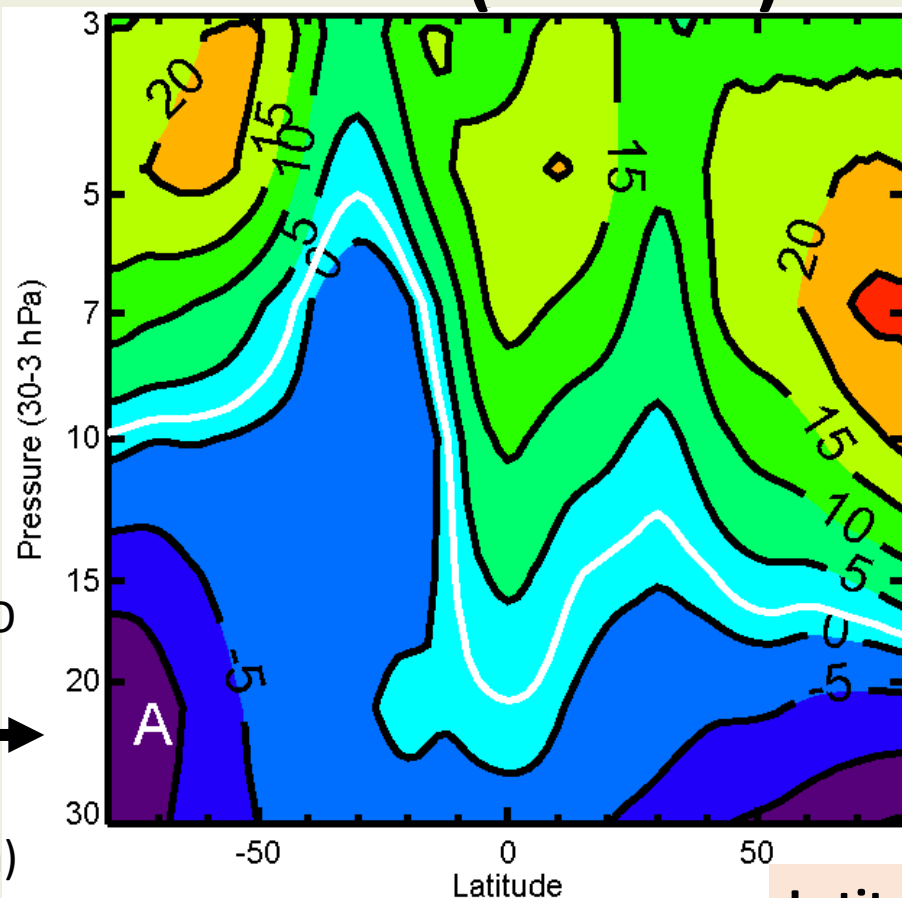
Quasi-Decadal (QD) N_2O Change over 2005-2021 (Percent change of 2013-2021 mean - 2005-2013 mean)

N_2O surface growth over this period = 2% (shown by white contour)

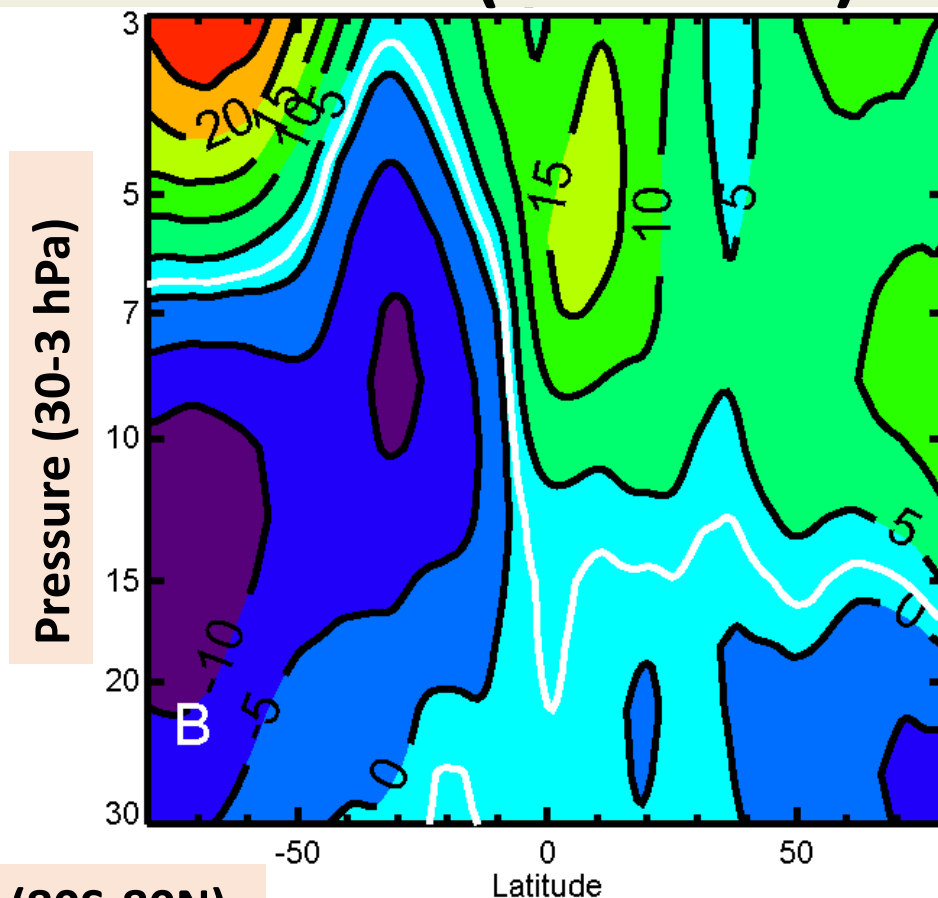
No drift in MLS N_2O 22 hPa and above

Known negative drift below 22 hPa)

Obs (MLS V5)

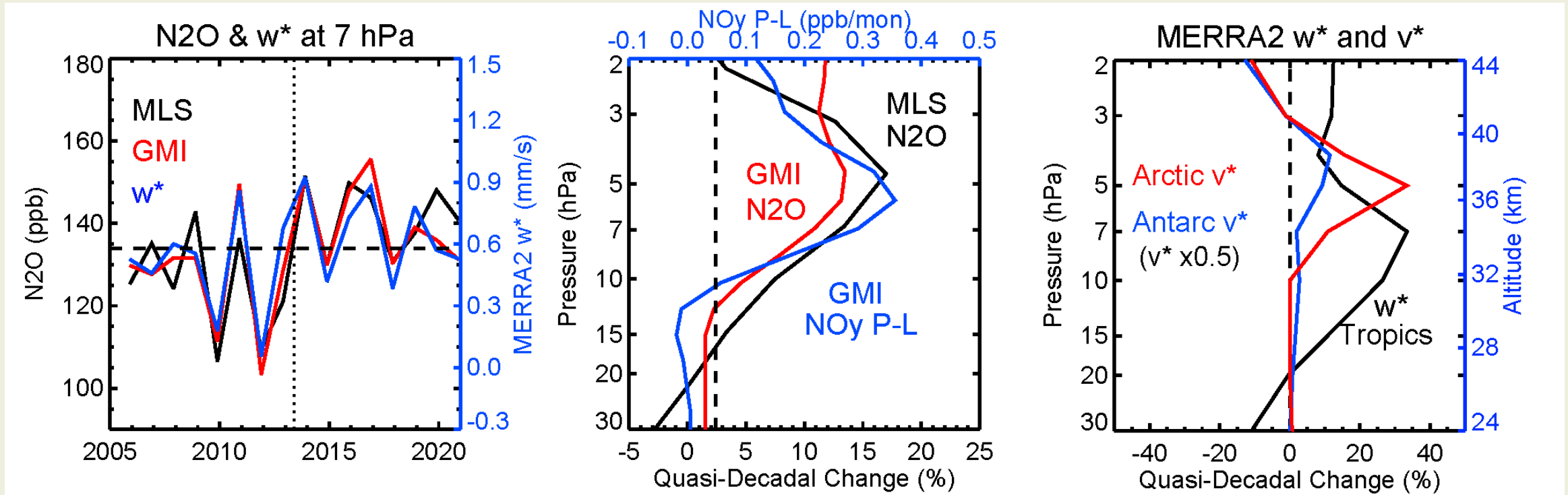


Model (GMI CTM)



Upper Stratospheric Circulation Change (w^* and v^*)

Explains why is N_2O so high



Tropical N_2O grows above 10 hPa, leading to increased odd nitrogen production.

Upwelling (w^*) increases above 20 hPa up to 30%

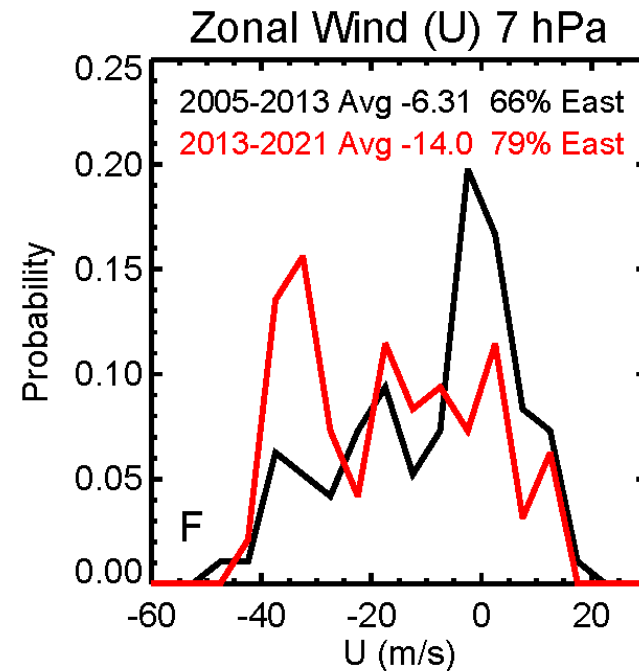
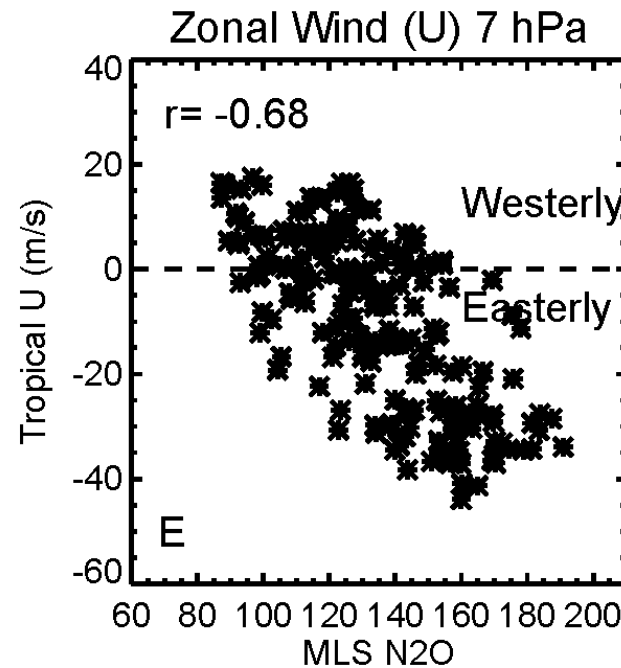
Arctic poleward transport (v^* 50-70°N) increases by 60% at 5 hPa

The Quasi-Biennial Oscillation (QBO) modulates tropical upwelling: stronger during Easterlies, weaker in Westerlies

MLS tropical monthly mean N_2O is positively correlated with w^* and anticorrelated with the zonal wind (i.e., QBO)

Stronger and more frequent QBO Easterlies 10-3 hPa increased mean tropical upwelling from 2013-2021

MLS and MERRA2 Monthly data from June 2005 to May 2021



2013-2021 was QBO-E
79% of the time

Two GMI Chemistry Transport Model (CTM) simulations with the same chemistry but different dynamics

GMI Chemistry Transport Model integrated with MERRA2:

- BASELINE: Time-varying MERRA2 Fields from June 2005-May 2021
- Fixed Dyn: June 2005-May 2007 Merra2 Fields (2 years) recycled until May 2021
- Both Simulations are forced with the same time-dependent source gases (N_2O , CFCs)

These start/end dates for the 2-year repeat chosen because of a 2-yr period QBO. This minimizes the transport adjustment when recycling.

O_3 loss differences between BASELINE and FIXED DYN show how dynamical changes led to chemical changes that affected O_3 .

Compare QD Change in O₃ Loss from NO_x and ClO_x (ppb O₃/mo) in BASELINE and Fixed Dyn Simulations

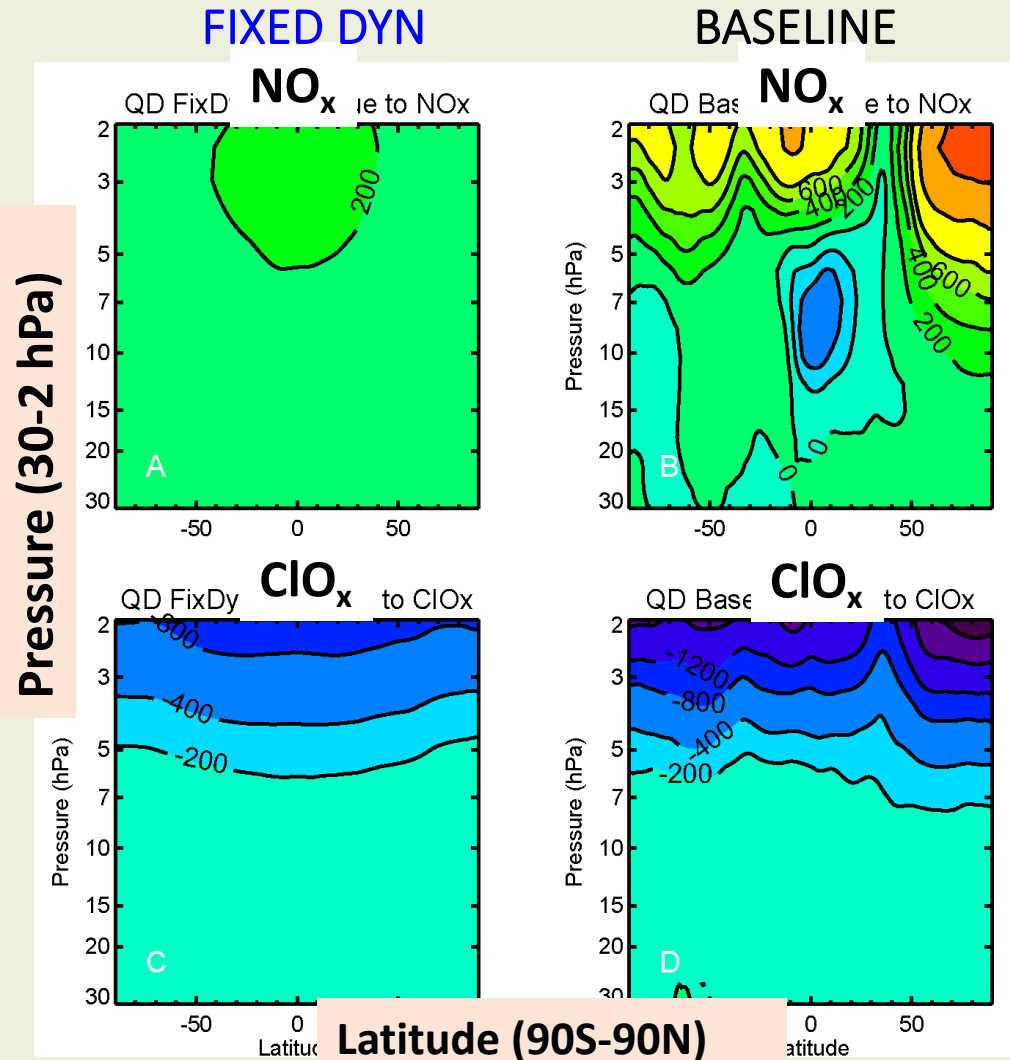
Fixed Dyn:

Hemispheric Symmetry

This is consistent with expectations from trends in tropospheric source gases (CFCs, N₂O)

Loss by NO_x increases above 5 hPa

Loss by ClO_x decreases with increasing altitude



BASELINE:

Hemispheric asymmetry

Greater loss by NO_x above 5 hPa, especially in the Arctic

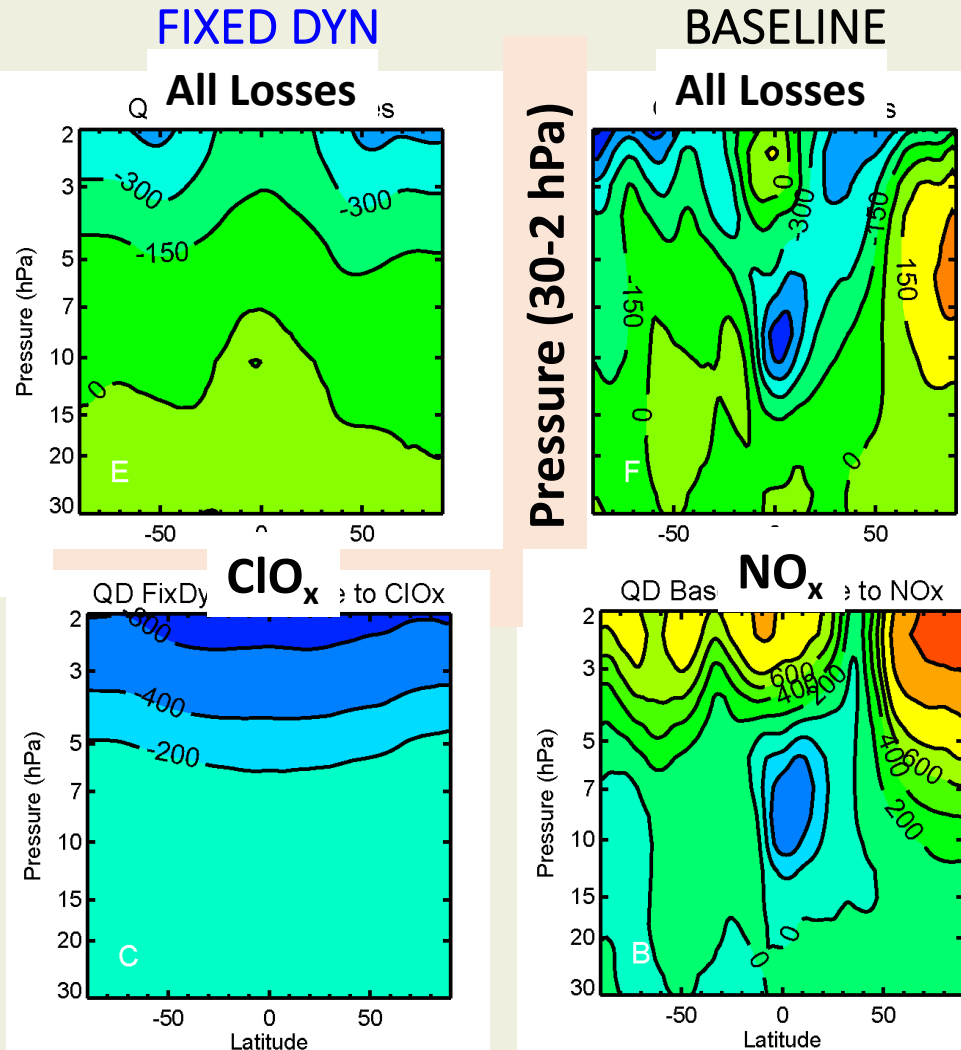
Less loss by NO_x in the tropical middle strat

Loss by ClO_x decreases even more than FixedDyn in the Arctic Upper Strat

The QD Change in the Sum of all O₃ Losses (NO_x, ClO_x, HO_x, etc.)

is driven by the NO_x Loss Changes in Baseline

Fixed Dyn: QD O₃ loss changes are dominated by Cl changes



BASELINE:

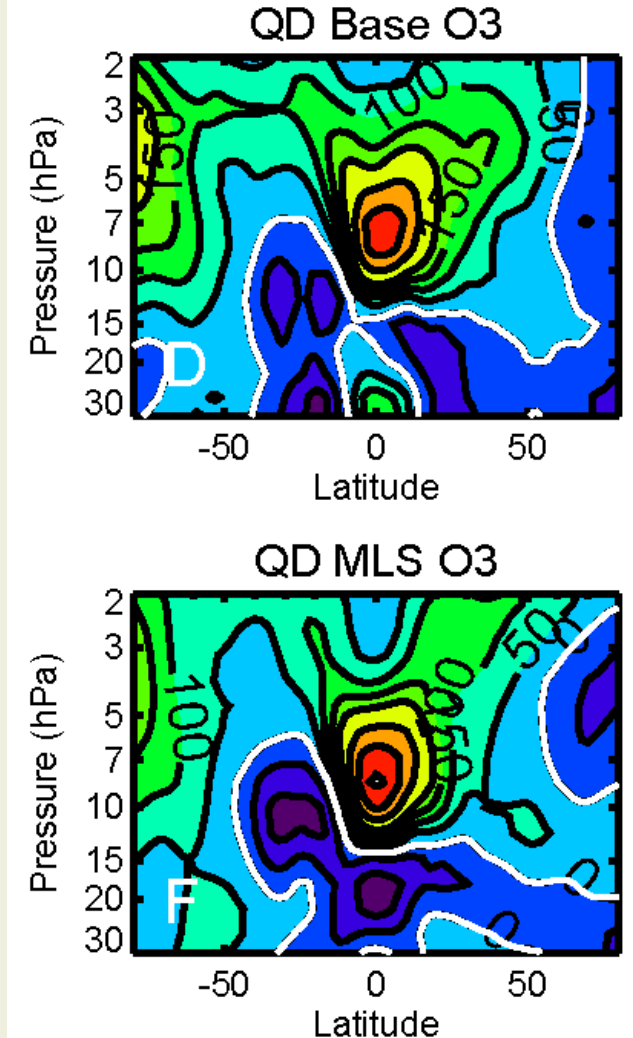
QD O₃ loss changes are dominated by NO_x changes. ClO_x changes contribute ~3 hPa and above.

And the denouement...BASELINE QD O₃ changes (ppb) look very much like observed O₃ changes 2005-2021

Arctic Upper Strat: O₃ decreases! It's not a lot (<200 ppb) but it's not increasing. Driven by NO_x

The Tropics: middle strat has a big O₃ increase because increased upwelling reduced NO_x and the loss it causes

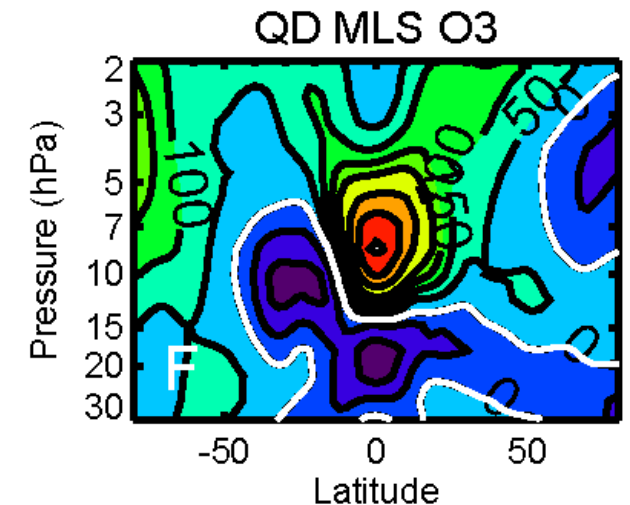
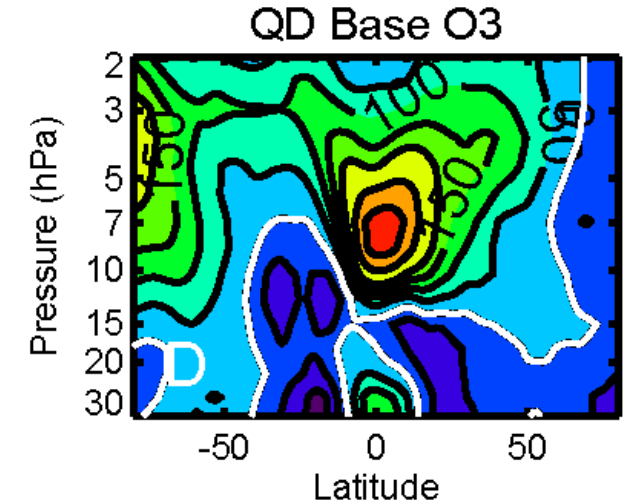
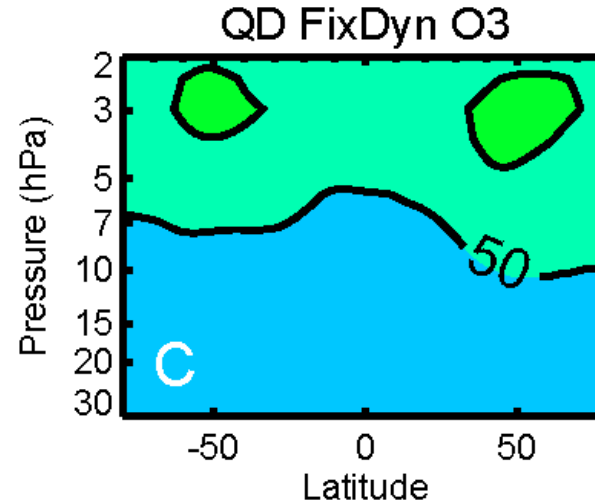
Antarctic Upper Strat: O₃ increases – reduced loss by ClO_x and less loss by NO_x (15-5 hPa)



And the denouement....Fixed Dyn QD O₃ changes look nothing like observed O₃ changes 2005-2021

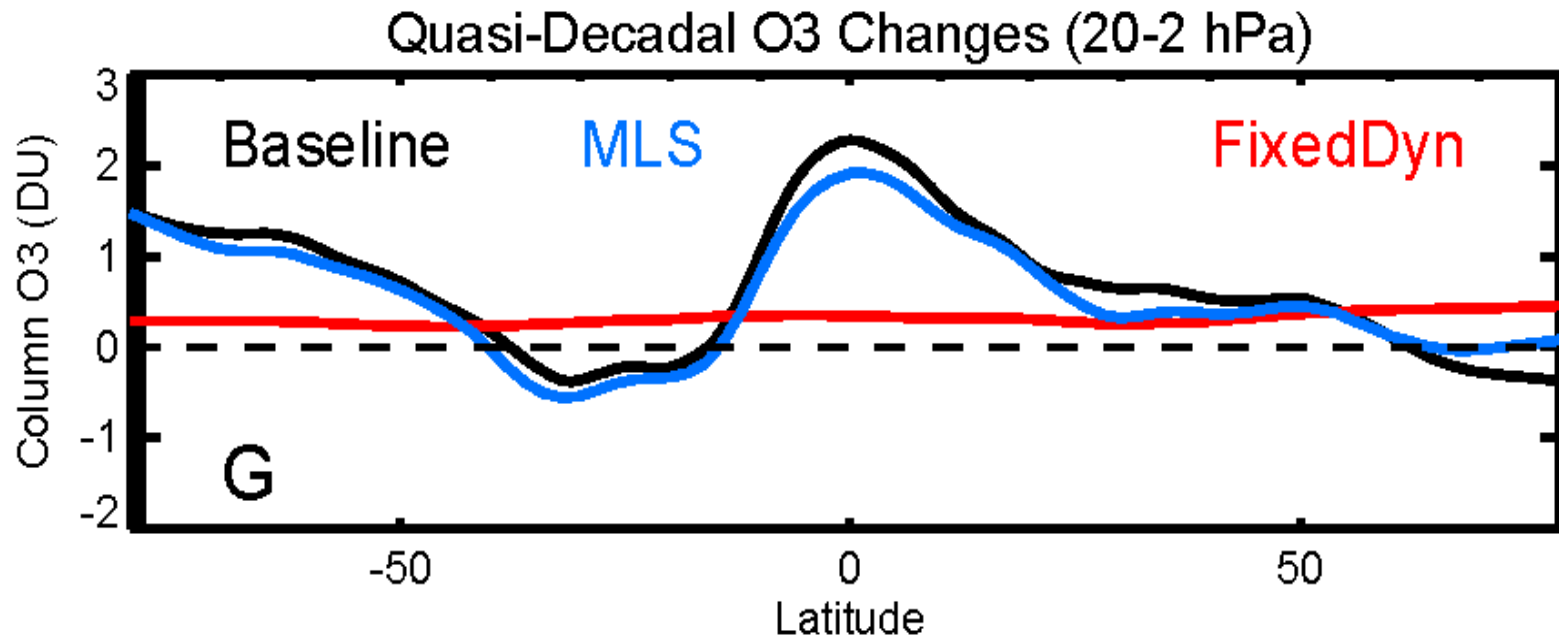
The Dynamically driven chemical changes are BASELINE – FixDyn O₃ (lower left panel).

Much of the observed O₃ changes 2005-2021 are caused by dynamically driven composition change!!



↑
O₃ Changes due to Source Gas Changes Only!!

Dynamically driven chemical changes from 20-2 hPa column Affect Total Column O₃ Trends



Check out the excellent agreement between MLS (blue) and the GMI CTM (black)

The 20-2 hPa column O₃ Quasi-Decadal changes caused by dynamics:

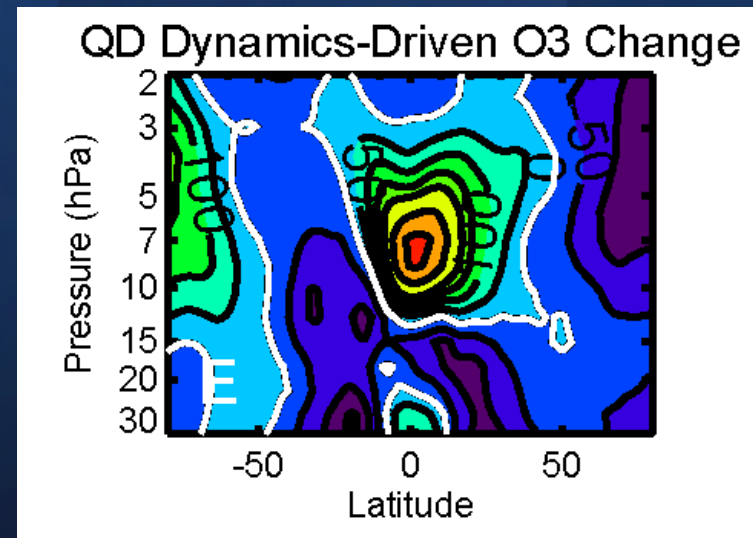
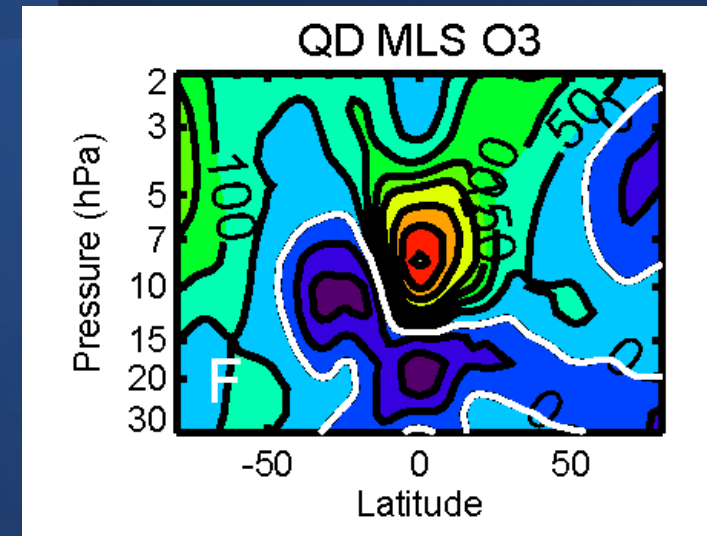
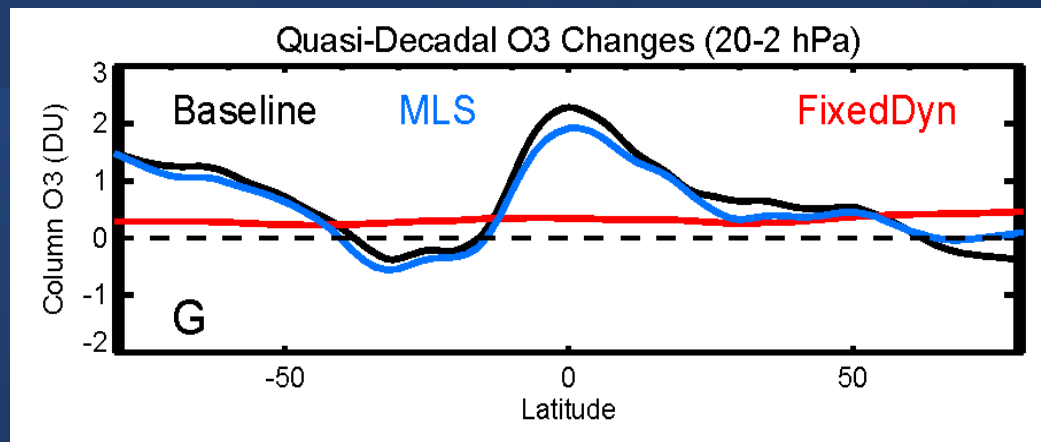
+2 DU in the tropics, -0.5 DU north of 60°N, and +1 DU in the Antarctic.

Questions/Thoughts...

Will QBO changes persist? Is this a trend or just variability?

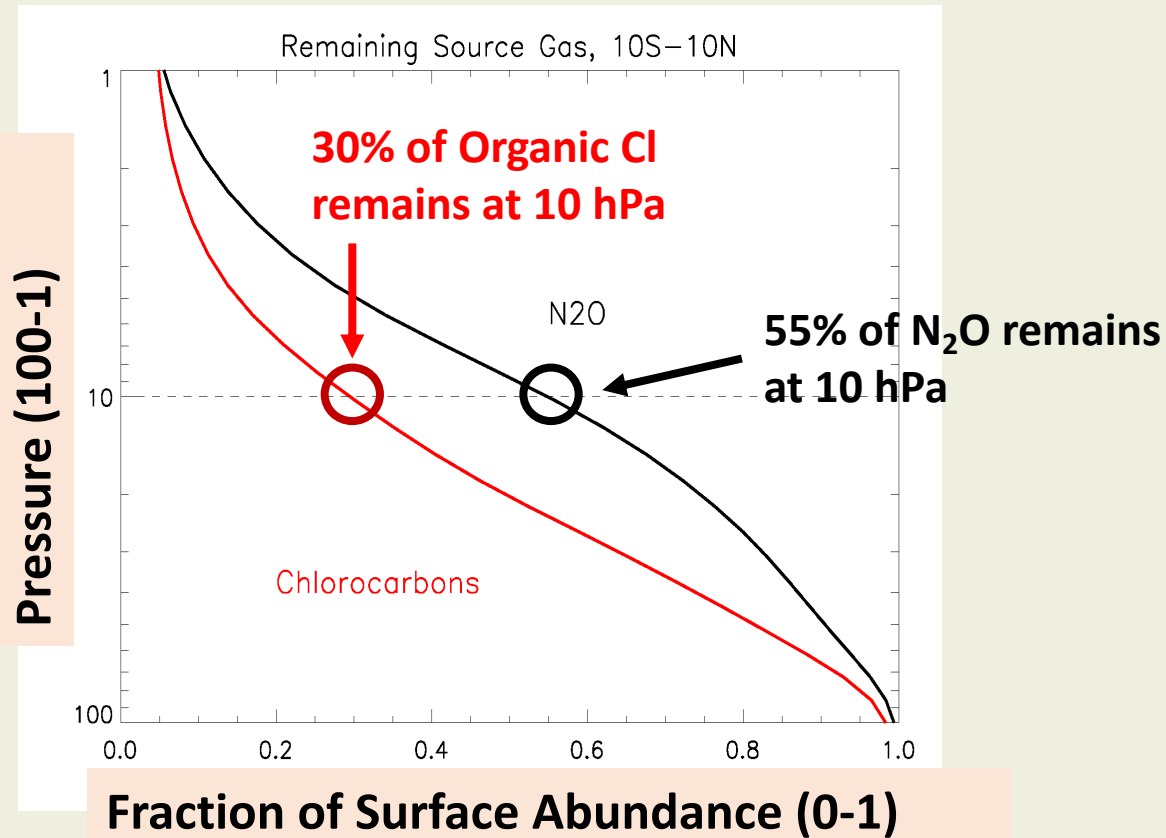
Chemistry Climate Models parameterize the QBO and can't respond physically to changes in forcing. They can't produce/predict this kind of effect on O_3 .

O_3 trend regressions fit the QBO with terms for the QBO 30 and 50 hPa (transport!). This can't regress changes in composition that affect O_3 chemistry



Please check out our 2022 GRL paper (Strahan et al.)

Why is there hemispheric asymmetry in Loss by NO_x but not the loss by ClO_x ?



Upwelling changes affect product gases (e.g., radicals) by changing source gas distributions

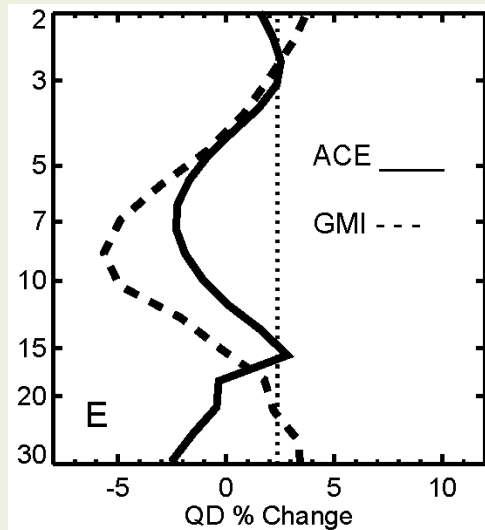
The upwelling changes are 10 hPa and above

The CFCs are mostly photolyzed below 10 hPa but most N_2O is still unreacted.

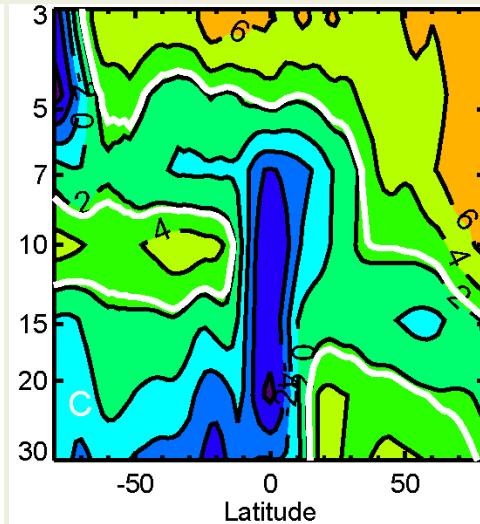
Chemical impacts of Circulation Change: Increased Net Production of NO_y is transported to the Arctic

Pressure (30-3 hPa)

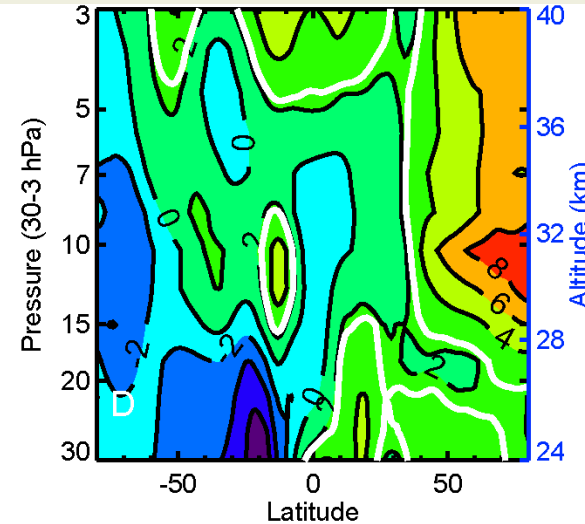
Tropical ACE NO_x



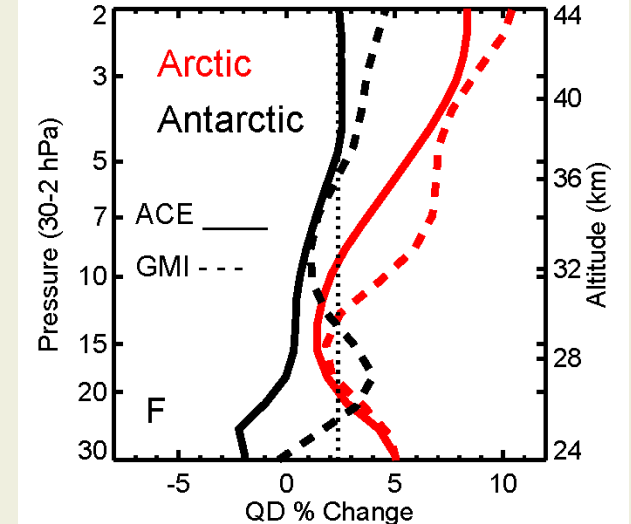
MLS HNO_3



GMI HNO_3



Polar ACE NO_x



In the tropics, N_2O is relatively long-lived below ~ 7 hPa, thus N_2O and NO_y are largely transport controlled.

N_2O and NO_y are anti-correlated

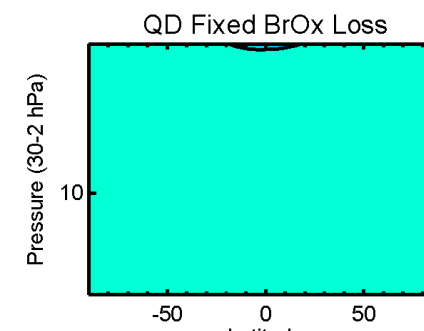
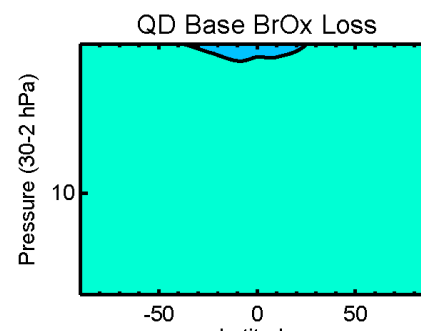
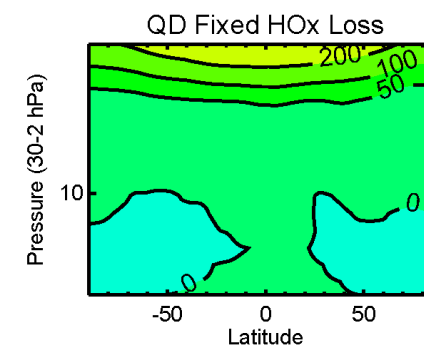
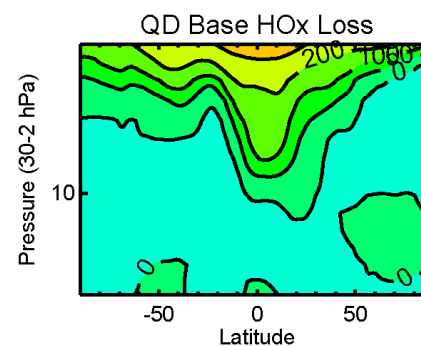
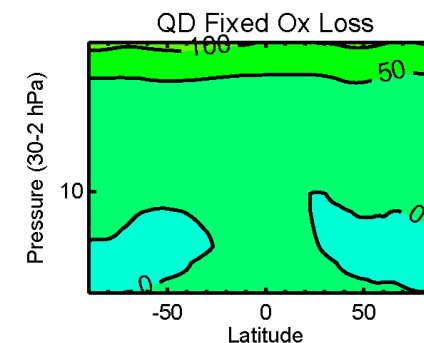
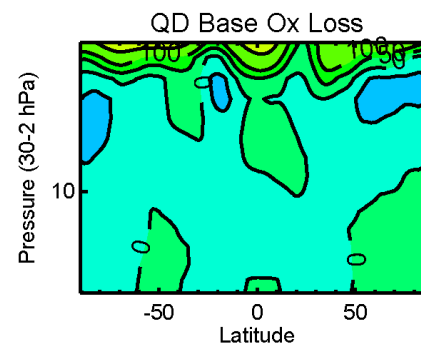
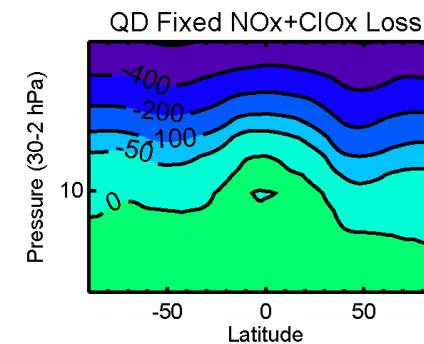
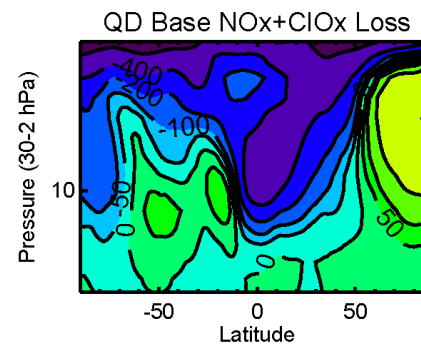
All O3 Loss Cycles (QD Change), Fixed Dyn (left) and Baseline (right)

NO_x + ClO_x

O_x

HO_x

BrO_x



And the denouement....Fixed Dyn QD O₃ changes look nothing like observed O₃ changes 2005-2021

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